

Ground motion model (GMM) for directional inelastic spectral displacements

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Introduction

- Intensity measure (IM) → links seismological conditions with engineering demands
- Ground motion models (GMMs) provide the probability distribution of an IM at a site, given underlying seismic hazard conditions
- Ground motions (GMs) can then be selected and scaled to match that IM distribution → and then use them for nonlinear response history analysis (NRHA) of structures
- Inelastic spectral displacement (Sd_i) can be an effective IM, under certain conditions
- Novel horizontal component definition for Sd_i : $RotD50$ and $RotD100$
 - ‘ $RotDnn$ ’ denotes the nn^{th} percentile of IM from all rotation angles sorted by amplitude
 - ‘D’ denotes that it’s dependent on the vibration period
- $Sd_{i, RotD100}/Sd_{i, RotD50}$ can be a more informative directionality measure, extended from $Sa_{RotD100}/Sa_{RotD50}$ which is the most common measure of GM directionality
- This $Sd_{i, RotD100}/Sd_{i, RotD50}$ can be also considered as a secondary IM

Luco and Cornell (2007)

Boore (2010)

Shahi and Baker (2014)



GMM for directional inelastic spectral displacements
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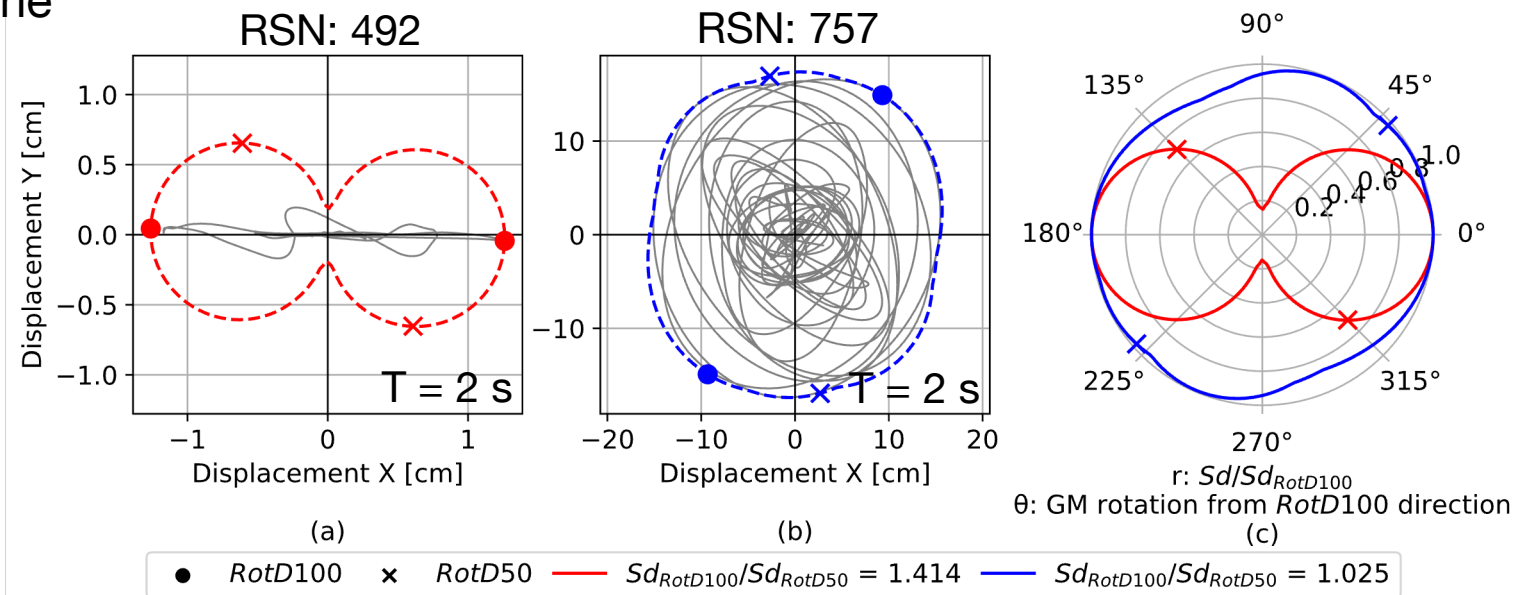
Ground motion directionality

- GMs include three translational and three rotational components (total: 6 components)
- Typically, rotational components are neglected, and vertical component receives much less attention than the horizontal ones
- Need to account for horizontal shaking in all orientations: when selecting and scaling GMs (with an IM that accounts for that), but also when applying them to structures
- Traditionally, the geometric mean of the IMs in each recorded direction was used, but orientation-dependent
- Commonly preferred orientation-independent definitions: *RotD50* and *RotD100*.

$$IM_{GM} = (IM_x \times IM_y)^{0.5}$$

$$IM_{RotD50} = \text{median}_{\theta}[IM(\theta)]$$

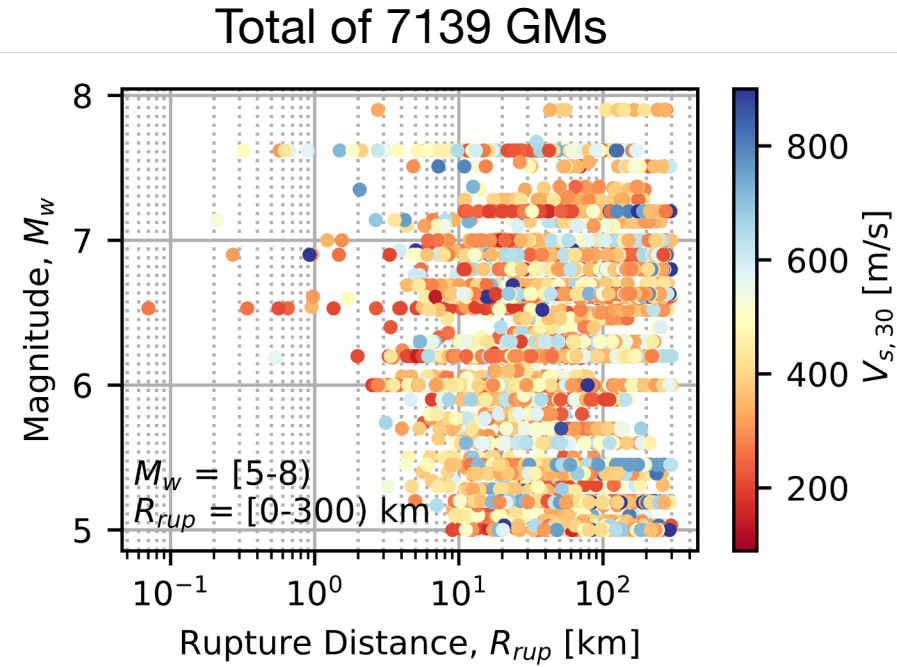
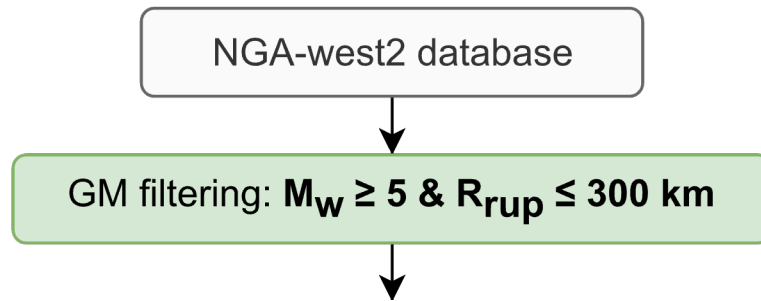
$$IM_{RotD100} = \max_{\theta}[IM(\theta)]$$



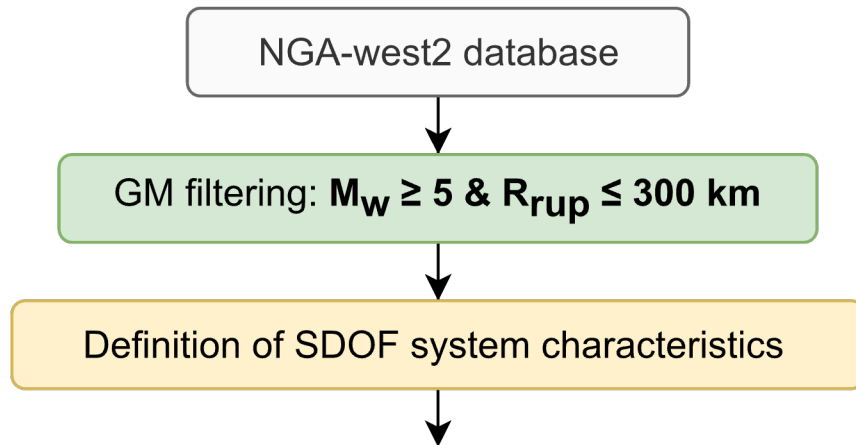
Baker, Bradley & Stafford (2021) – Section 4.2.6



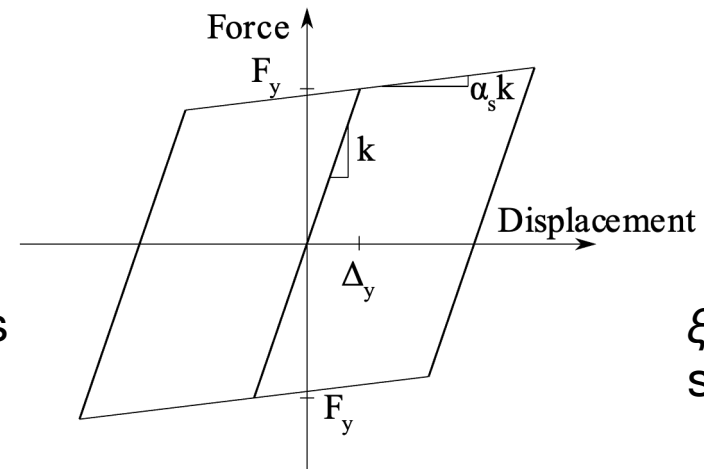
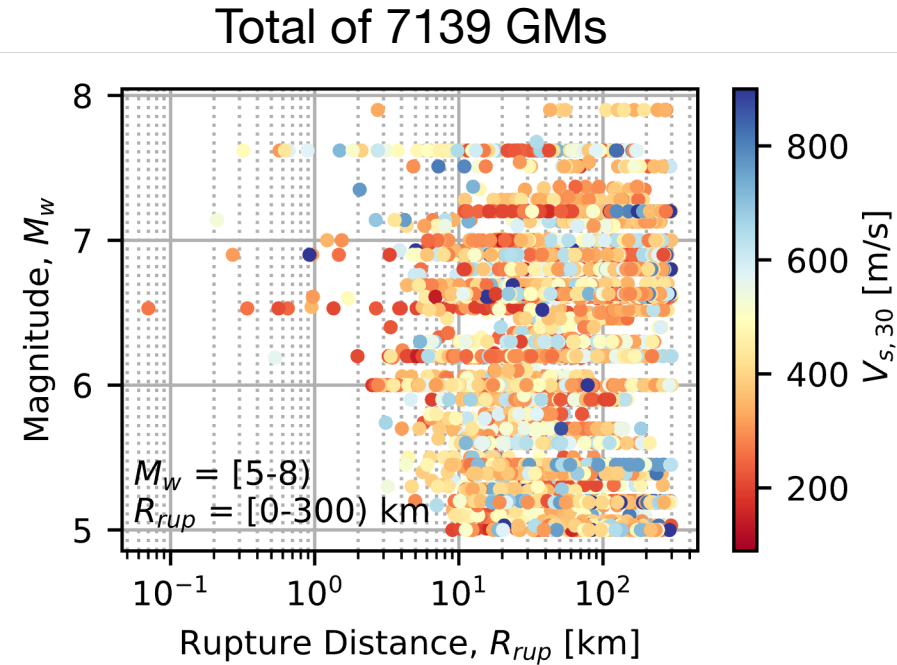
Methodology



Methodology



$T = [0.04, 0.06, 0.1, 0.2, 0.3, 0.5, 0.75, 1, 1.5, 2, 3, 4, 5] \text{ s}$
 $R = [1.5, 2, 3, 4, 6]$

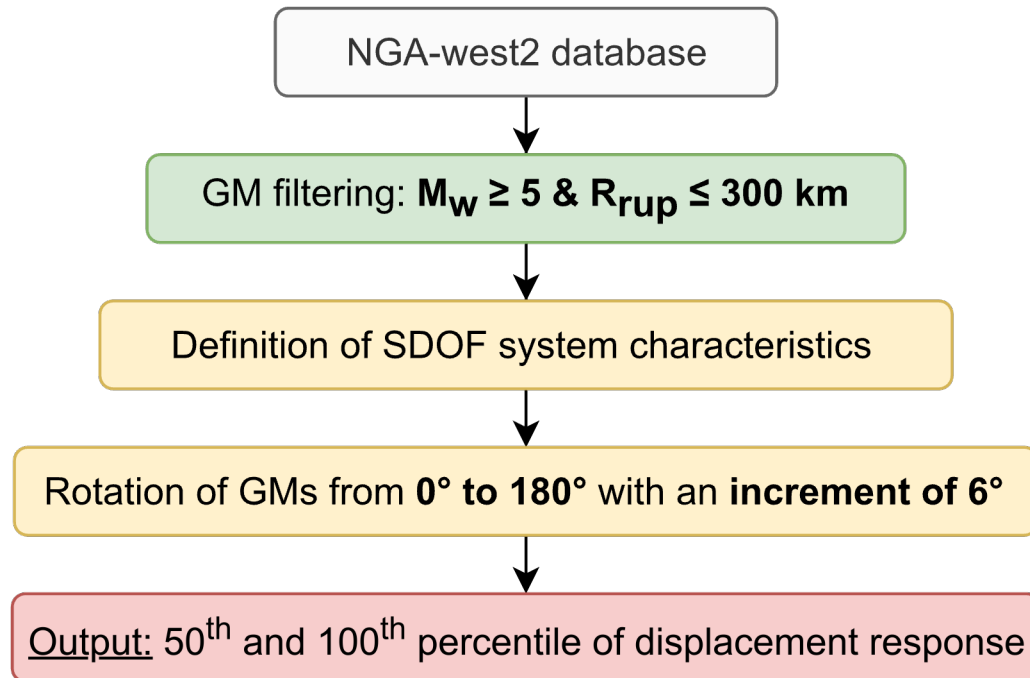


$$F_y = \frac{m \times S a_{RotD100}}{R}$$

$$a_s = 3 \%$$

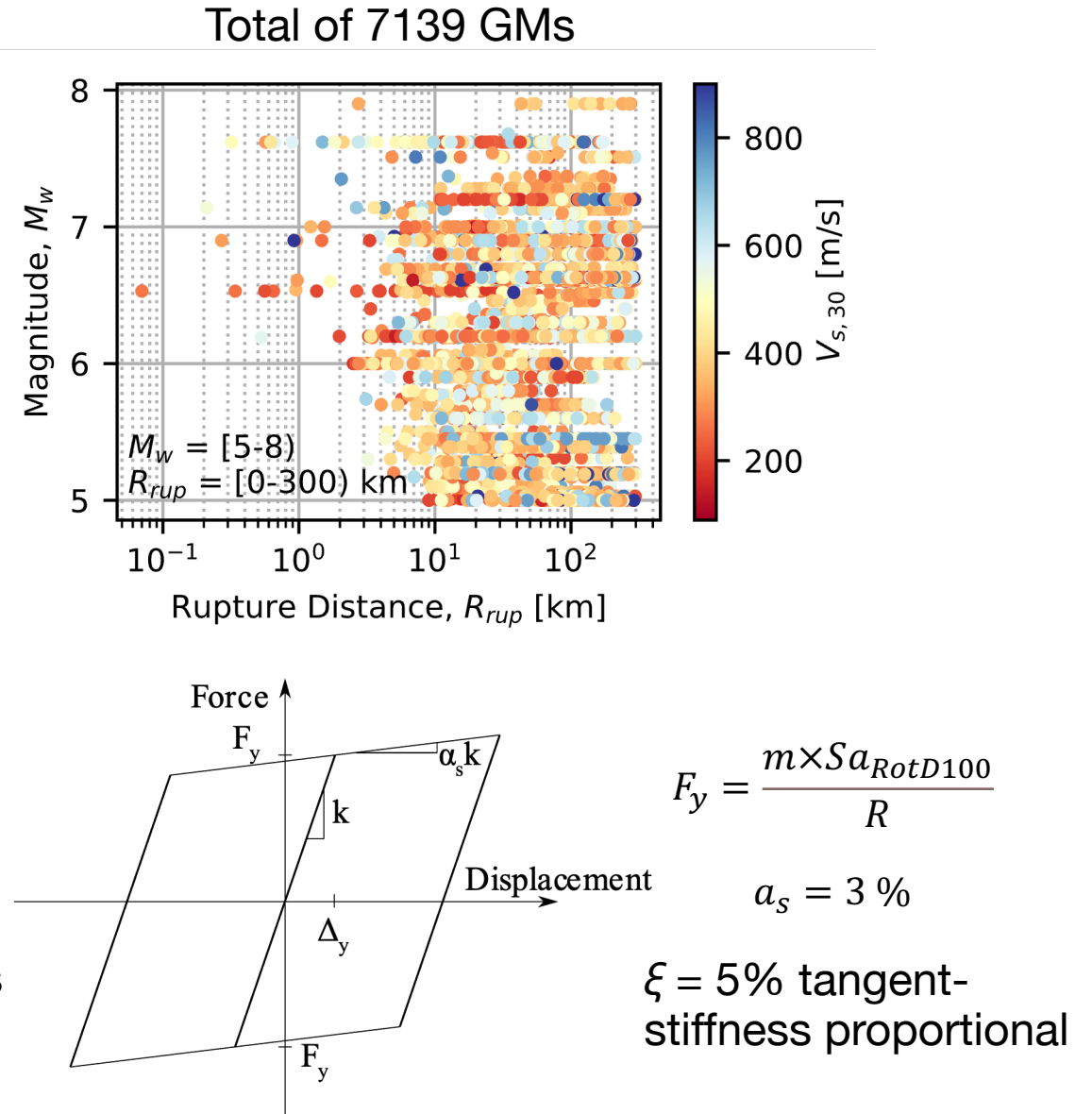
$\xi = 5\%$ tangent-stiffness proportional

Methodology

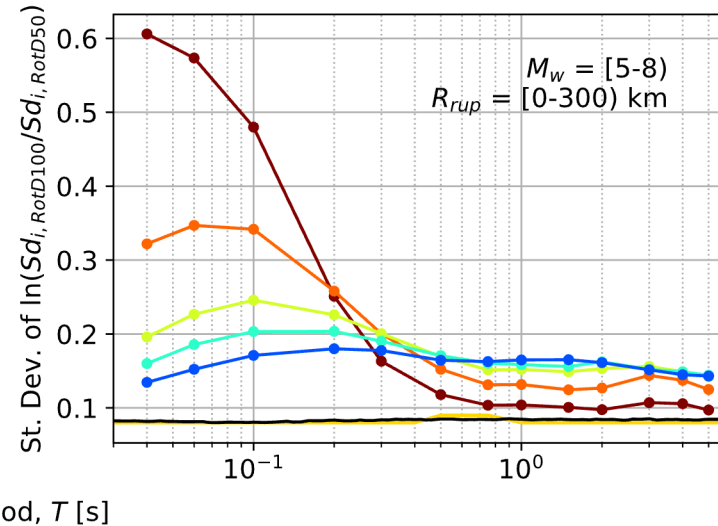
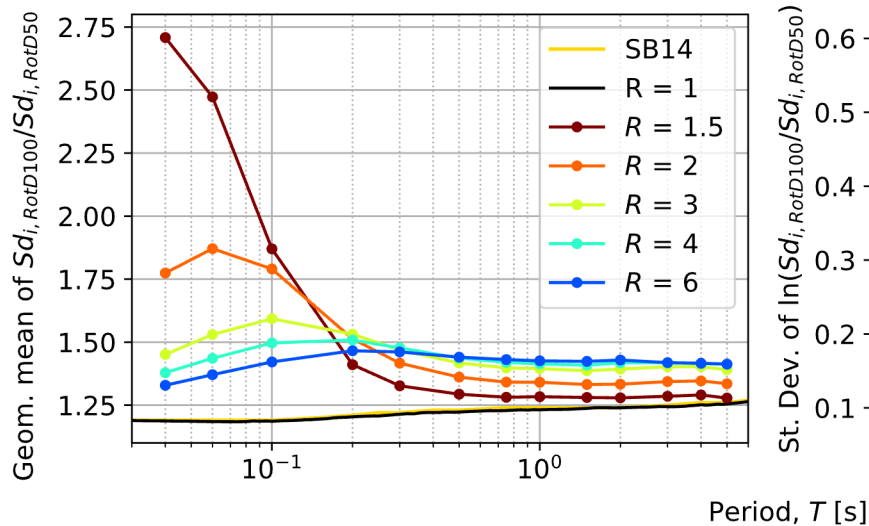


$T = [0.04, 0.06, 0.1, 0.2, 0.3, 0.5, 0.75, 1, 1.5, 2, 3, 4, 5]$ s

$R = [1.5, 2, 3, 4, 6]$



Directionality model for inelastic displacements



| R = 2 | | | | | |
|-------|--------|-----------|----------|-------------------------------------|--------------------------------|
| T [s] | τ | φ | σ | $\ln(Sd_{i,RotD100}/Sd_{i,RotD50})$ | $Sd_{i,RotD100}/Sd_{i,RotD50}$ |
| 0.04 | 0.099 | 0.306 | 0.321 | 0.651 | 1.917 |
| 0.06 | 0.079 | 0.334 | 0.344 | 0.679 | 1.973 |
| 0.1 | 0.044 | 0.333 | 0.336 | 0.581 | 1.788 |
| 0.2 | 0.040 | 0.249 | 0.252 | 0.397 | 1.487 |
| 0.3 | 0.023 | 0.196 | 0.197 | 0.339 | 1.404 |
| 0.5 | 0.012 | 0.150 | 0.150 | 0.307 | 1.360 |
| 0.75 | 0.006 | 0.130 | 0.130 | 0.295 | 1.343 |
| 1 | 0.007 | 0.130 | 0.130 | 0.292 | 1.339 |
| 1.5 | 0.011 | 0.122 | 0.123 | 0.284 | 1.328 |
| 2 | 0.009 | 0.125 | 0.125 | 0.287 | 1.333 |
| 3 | 0.021 | 0.137 | 0.138 | 0.291 | 1.338 |
| 4 | 0.019 | 0.130 | 0.131 | 0.292 | 1.339 |
| 5 | 0.015 | 0.121 | 0.122 | 0.287 | 1.332 |

Shahi and Baker (2014) \longrightarrow Directionality model for $Sa_{RotD100}/Sa_{RotD50}$

Aristeidou, Tarbali, and O'Reilly (2023) \longrightarrow Directionality model for $Sd_{i,RotD100}/Sd_{i,RotD50}$

Functional form of GMM

$$\ln Y_{i,j} = a + F_M + F_D + F_{sof} + F_s + F_{basin} + \eta_i + \varepsilon_{i,j}$$

$Sd_{i,RotD50}$ or $Sd_{i,RotD100}$

- a:** Model scaling coefficient
- F_M:** Magnitude scaling term
- F_D:** Distance attenuation term
- F_{sof}:** Style of faulting term
- F_s:** Site amplification term
- F_{basin}:** Basin effects correction term
- η_i:** inter-event residual
- ε_{i,j}:** intra-event residual

Predictor seismological parameters:

M_w: Moment magnitude

R_{rup}: Rupture distance

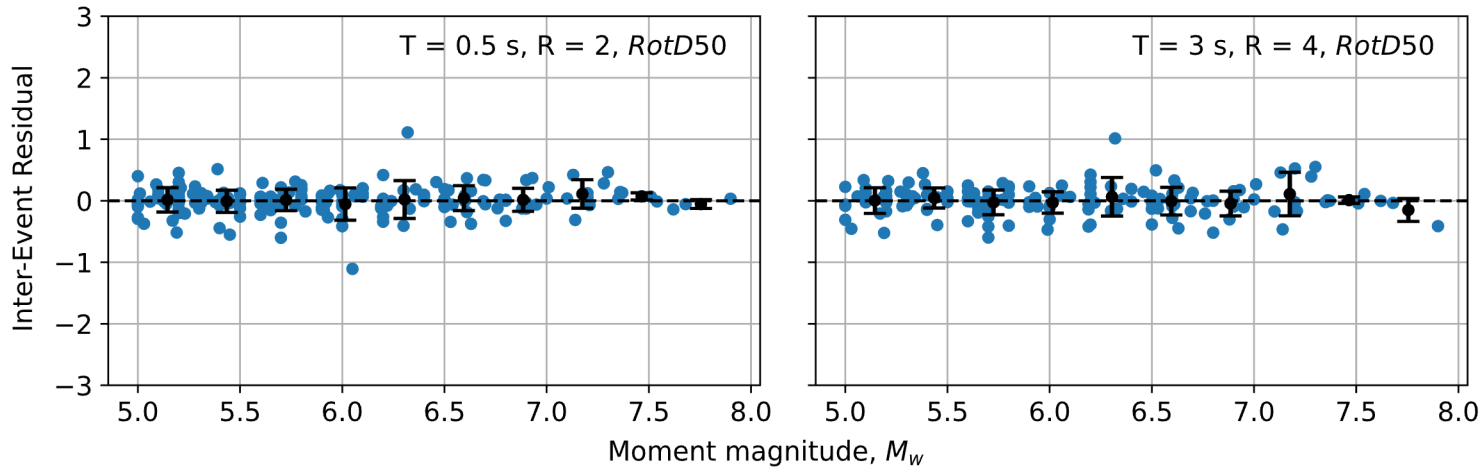
Fault mechanism: Discretised into 3 faulting styles: strike-slip, normal and thrust fault

V_{s,30}: Time-averaged soil shear-wave velocity to 30 m depth

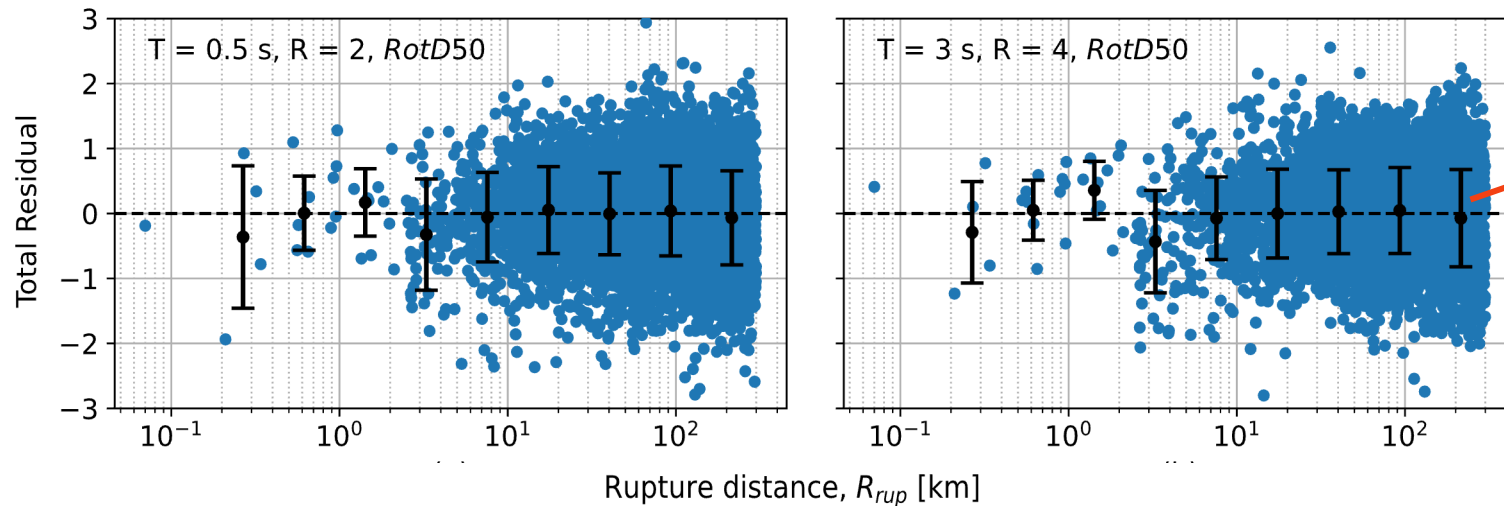
Z_{2.5}: Depth to the 2.5 km/s shear-wave velocity horizon (basin proxy)

18 model coefficients calibrated for each elastic vibration period, **T**, and strength ratio, **R**

GMM performance



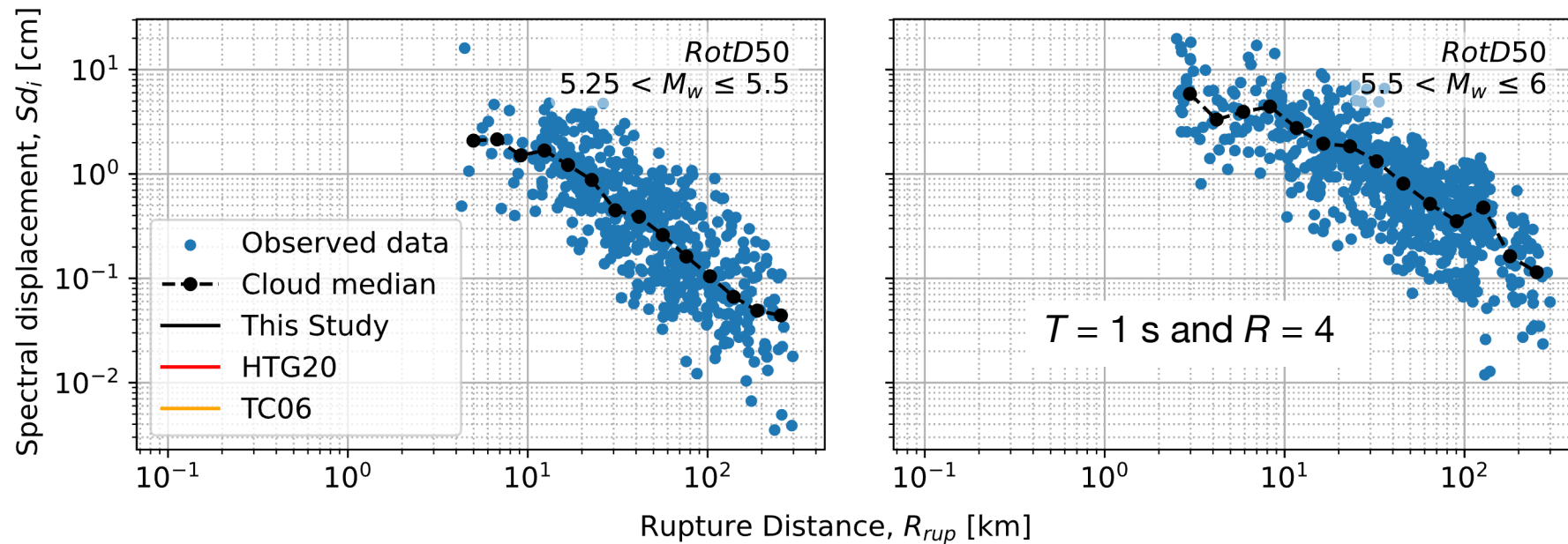
- Residual definition: difference between ‘observed empirical data’ and model predictions
- No apparent bias is present
- $R^2 \approx 0.8$ for most T and R cases



Binned mean residual values
 ± 1 standard deviation

Comparison with other studies

- Only a few available models that have strength ratio, R , as input. A few of them have ductility demand, μ , or strength coefficient, C_y
- Two models from the literature were compared herein
- Median prediction of the proposed GMM matches well the cloud median



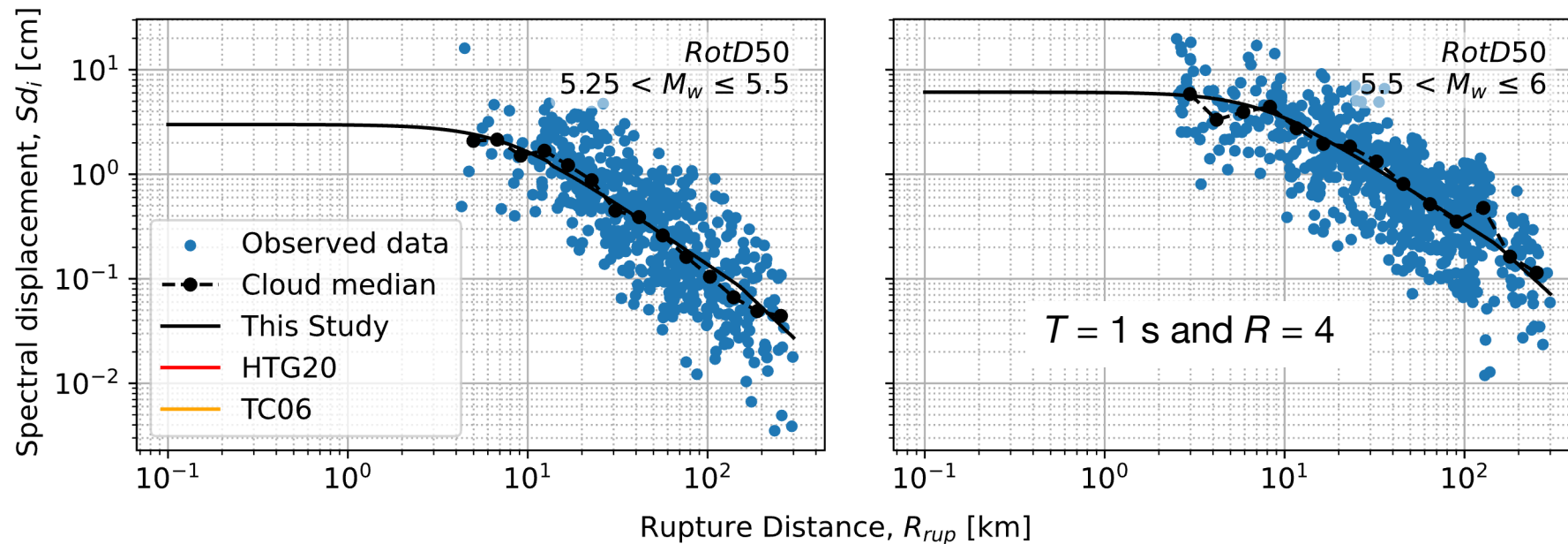
Huang, Tarbali and Galasso (2020): HTG20

Tothong and Cornell (2006): TC06



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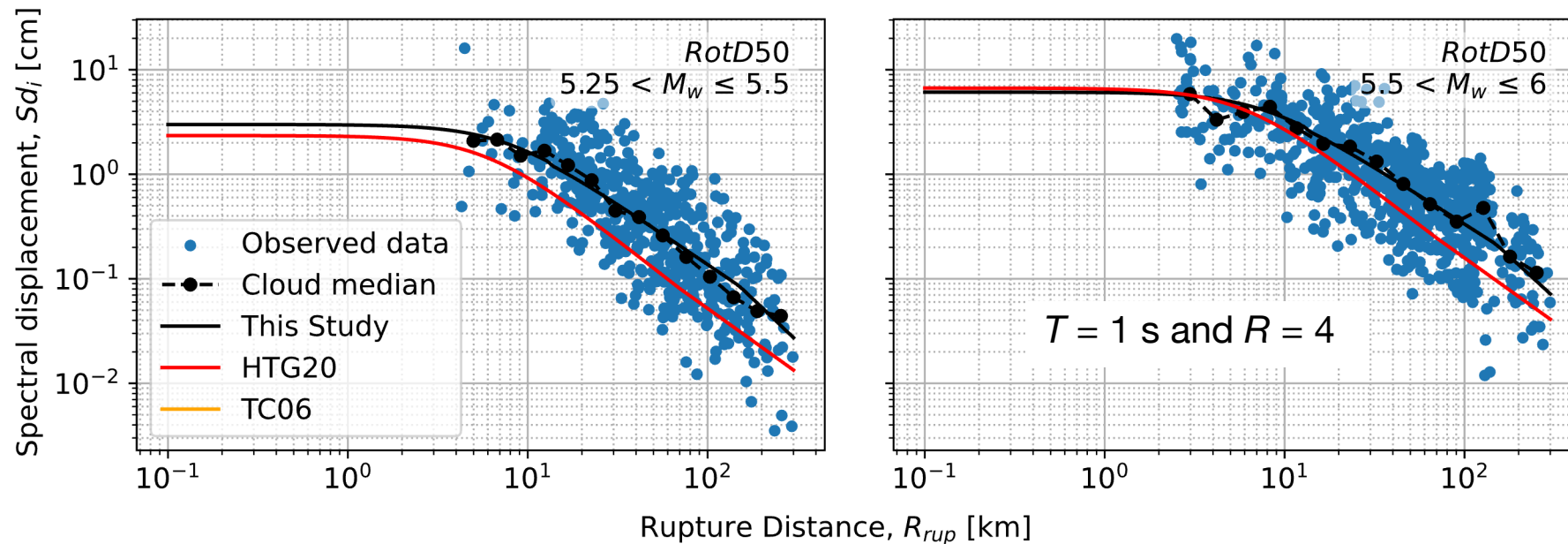
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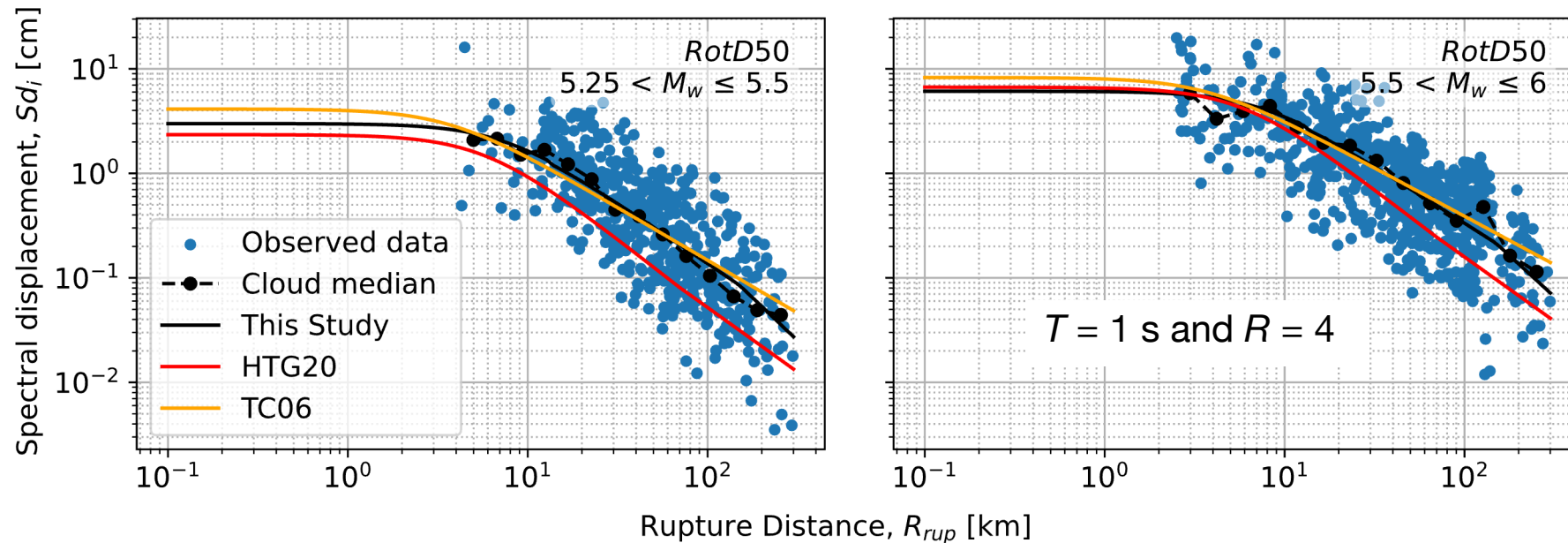
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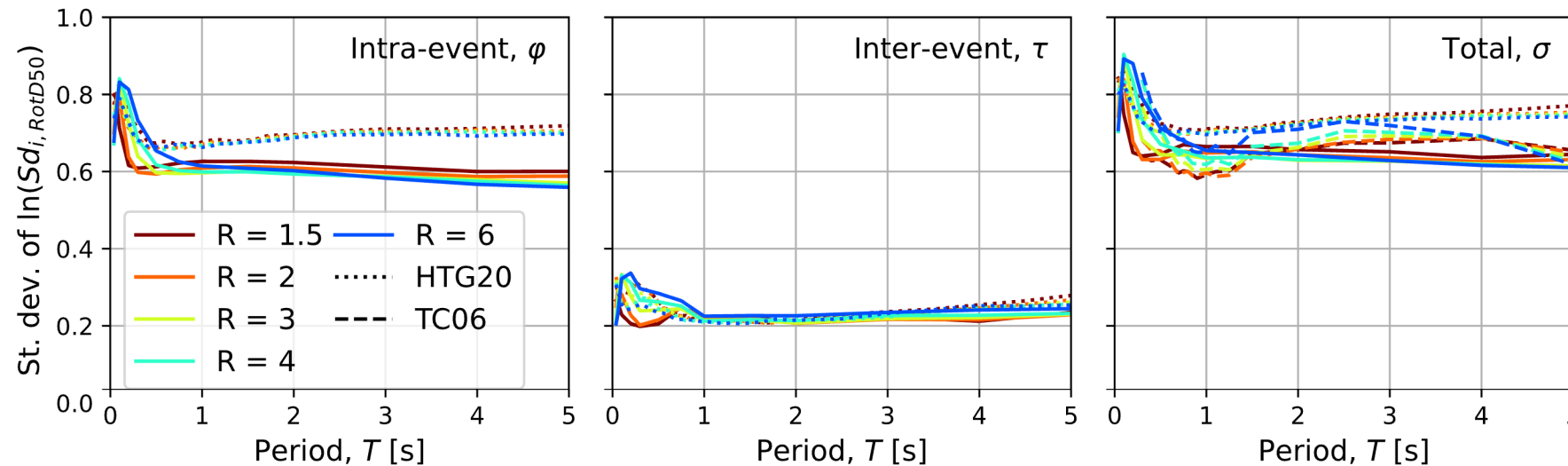
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Variability in the GMM

- The proposed model gives lower standard deviations for most periods when compared to TC06 and HTG20
- *RotD50* component slightly reduces the dispersion in comparison to the arbitrary component used by TC06 and to the geometric mean used by HTG20
- HTG20: difference mainly due to intra-event, which is a product of considering spatial correlation



Inter- and intra-event residuals mutually independent

$$\sigma = \sqrt{\tau^2 + \varphi^2}$$

Beyer and Bommer (2006)

Jayaram and Baker (2010)



Conclusions

- GMM developed to estimate the *RotD50* and *RotD100* horizontal component of Sd_i from shallow-crustal earthquakes
 - Used substantially large dataset of GMs from NGA-West2 database
 - Does not require any auxiliary elastic GMM to predict the median and dispersion of inelastic displacements
 - Range of applicability: $5 < M_w \leq 8$; $0 < R_{rup} \leq 300$ km; $90 \leq V_{s,30} \leq 1300$ m/s; $0.04 \leq T \leq 5$ s; $1 \leq R \leq 6$; tectonically active shallow crustal regions
 - Model exhibits good performance and reasonably low dispersions, compared to similar models available in literature, and they are not sensitive to the level of non-linear demand
 - Proposed directionality models based on Sd_i , given in the journal paper, can be used
 - Directionality can be also estimated from the GMM itself, using the different available horizontal component definitions
-
- *Aristeidou, S., K. Tarbali, and G. J. O'Reilly. 2023. "A ground motion model for orientation-independent inelastic spectral displacements from shallow crustal earthquakes." Earthq. Spectra, 0 (0): 1–23. <https://doi.org/10.1177/87552930231180228>.*
 - *Aristeidou, S., G. J. O'Reilly. 2023. "Exploring the use of orientation-independent inelastic spectral displacements in the seismic assessment of bridges." Under review.*



Thank you for your attention!



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